obtain reliable means for each month. The relations, if any exist, between the daily variability of temperature and the mean monthly temperature have not been very thoroughly investigated. For instance, it is possible that in extremely cold months the variability of temperature is greater, and in warm months less, than the average.

Using data compiled by Wahlen from long records of Russian stations, the author has undertaken this investigation. The Russian stations used are Warsaw (101 years), Port Baltic (44 years), Archangel (68 years), Astrakhan (44 years), Katharinenburg (52 years). Barnaul (45 years), and Yakutsk (37 years). In addition, data for Vienna (75 years) and for Pola (40 years) com-

piled by von Hann, were used.

The results of the investigation were three fold: For Pola, which typifies the climate influenced by the ocean, there was no relation between the variability and the departure from the monthly mean. Vienna, Warsaw, Port Baltic, Archangel, and Astrakhan fall into a second group, in which the following relation seems to hold: In the months colder than normal the variability of daily temperatures is greater than the average; in months warmer than normal, it is less than the average. This region seems to form a transition between oceanic and continental climates. At Barnaul and Katharinenburg in west Siberia, this rule does not hold, and at Yakutsk in east Siberia the relation is reversed; i. e., in winter months colder than normal there is a small variability of temperature, and in those warmer than normal the variability is large.

The proof that these relations are not accidental is to be found in the smooth progression of values from Pola to Yakutsk. The relation is weakest in regions of most intense cold. Great variability of temperature implies cold and warm waves following one another rapidly. But in Yakutsk the lower the monthly mean lies the more uniform the temperature, because of fewer warm

waves.1—C. L. M.

THE LAWS OF APPROACH TO THE GEOSTROPHIC WIND.

By F. J. W. WHIPPLE.

[Abstracted from Quar. Jour. Roy. Mct'l Soc., 46, 39-53, Jan., 1920.2]

In the steady state, the forces at each level in the atmosphere due to rotation, gradient, and internal friction (eddy viscosity) must balance; in the layer next the surface, the additional drag due to friction is practically equal and opposite to the drag due to the eddies, since it may be shown that the combined effect of pressure and rotation is only one-thirtieth that of one of the other two forces, and hence may be ignored. The actual wind at any level is considered as the vector sum of the gradient wind and a fictitious "relative wind." In the case of straight, parallel, motionless isobars, with uniform distribution of density, gradient, and eddy motion throughout the height considered, the above conditions lead to three laws: I. The relative wind falls off with increase of height according to the exponential law $R = R_0 e^{-(Bz-z_0)}$, where R is the relative wind at the height z, and R_0 that at height z_0 ; B is a constant, $R_0 = \omega \sin \lambda \ (\rho/\mu)$. II. The direction of the relative wind turns uniformly with increase of height, the circular measure of its azimuth being given by $\phi - \phi_0 = B (z - z_0)$.

III. The angle between the actual wind near the surface and the relative wind there is 135°. The velocity curve traced out by the lines, drawn from a fixed origin, representing the relative winds at all heights, is the equiangular spiral of Ekman, the constant angle between the radius vector and the tangent (the latter representing transfer of momentum by the eddies) being 45°. The Law of Dynamic Similarity then leads to Taylor's relation, $V_0 = G(\cos \alpha - \sin \alpha)$.

The first two laws also hold when G changes with height: $G = G_0 + Az$; then $e^{-3i\pi/4}$ $R_0 = a(V_0 - U)$, where $U = A/aB\sqrt{2}$, and $a = K_{\rho}V_{0}/\mu B\sqrt{2}$, a generalization of Law III. With V_{0} , and U known, G_{0} may be computed; and with the law of change of G with height known, V at any level may be computed. The theory is applied to the case where observation shows east winds at the surface to be replaced by west winds at 2,000 meters, so that the change of gradient wind with increasing height is equivalent to the addition of a westerly component; then for surface winds of equal strength, it is found that the backing of the surface wind is least for south winds, greatest for north winds; the ratio of surface wind to gradient wind is least for southeast gradient wind and greatest for northwest. This shows the tendency for the surface wind to be governed by the geostrophic wind aloft rather than the geostrophic wind at the surface.

The theory is applied in detail to Dobson's Salisbury Plain observations, and again the fact is theoretically confirmed that the gradient speed is attained at about 300 meters, while the gradient direction is not attained until a much greater height. Comparisons of theory with observation are rendered difficult because of the influence of the anemometer exposures, but such comparisons as have been made, especially with the data of T. Koraen, tend to confirm the theory—south winds being much lighter in comparison with gradient winds

than those from northwest.

(f. Monthly Weather Review: July, 1915, 43: 315-316; Sept., 1917, 45: 455; Jan., 1918, 46: 22; May, 1918, 46: 211; Oct., 1919, 47: 703-707.—E. W. W.

TECTONIC EARTHQUAKES AND VARIATIONS OF LATITUDE.

By G. ZEIL.

[Abstracted from Comptes Rendus, Paris Academy of Sciences, t. 171, pp. 311-313, Aug. 2, 1920.]

Omori, Milne, and Cancani found that destructive earthquakes took place at or near times of maximum or minimum variation of latitude. Brillouin concluded that variations of latitude were caused by sudden internal movements, of seismic origin; Montessus de Ballore's former doubts about this would not now hold, since he has come to believe that earthquakes and epeirogenic movements are one and the same. (C. R., 158, 1833,

The present author states that on a parallel of latitude, the various molecules of the lithosphere are in equilibrium two by two, at the opposite extremities of diameters of the circle; if one of these molecules is displaced along the diameter, a seesaw movement of the diameter will take place such that the displaced extremity will be above the normal level of rotation if it approaches the center, and below if it recedes from the center. On the basis of his former studies (see Monthly WEATHER REVIEW, 48, 356, June, 1920), Zeil then

¹ The colder the air, the stronger the disturbance necessary to blow it away. But in the coldest weather in continental interiors, such as central siberia, the winds are usually weak. The warm winds of winter may glide over the cold surface layers for many hours and may not reach the surface at all before they cease to blow.—EDITOR.

² Cf. Science Abstracts, Apr. 30, 1920, p. 518.

contends that in this manner centrifugal earthquakes will cause the polar axis to incline toward the readjusted arch and increase the latitude of the disturbed area; while centripetal quakes will cause the opposite. Antagonistic quakes will have little or no effect. It has been shown that the Japanese quakes are of two typesthose of submarine origin, essentially centripetal, and those of continental origin, essentially centrifugal—and their different methods of action explain why the quakes occur either at maximum or minimum variation of latitude.—E. W. W.

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